

Impact of surface emissions on the zonal variability of tropical tropospheric ozone and carbon monoxide for November 2004

Kevin Bowman, Helen Worden, Line Jourdain,
John Worden, Greg Osterman, Susan Kulawik
Jet Propulsion Laboratory, California Institute of Technology

Dylan Jones

University of Toronto

Jennifer Logan, Folkert Boersma

Harvard University

Juyun Warner

University of Maryland, Baltimore County

Henk Eskes, Pepijn Veefkind, and Ronald van der A
KNMI

Surface emissions and ozone

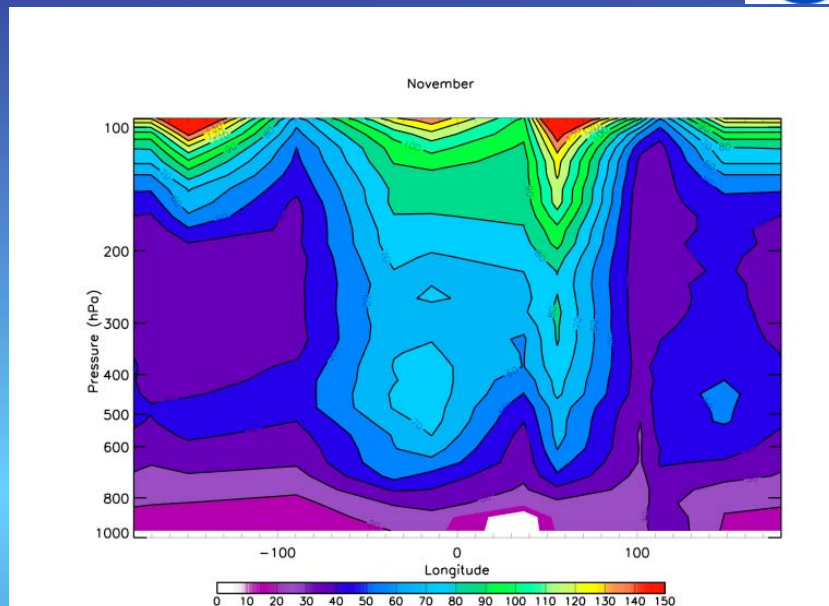
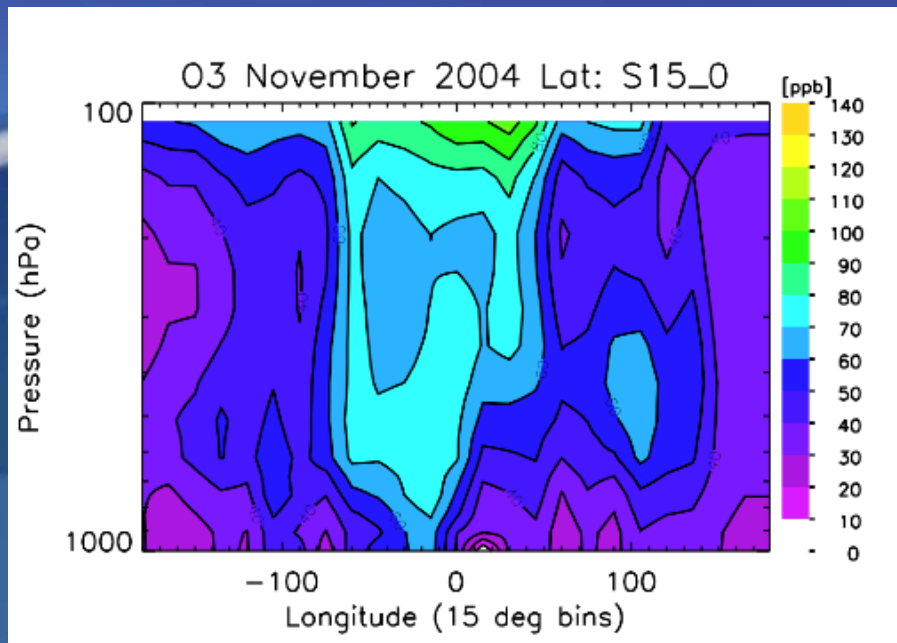
- Emissions from biomass burning and biofuel are the largest anthropogenic contributor of ozone precursors in the southern hemisphere
- The scale, distribution, and composition of these emissions are subject to large uncertainties
- Ascribing ozone formation to these emissions is complicated by significant natural sources of ozone and transport
- Satellite observations--in particular TES--can be combined with chemical transport models via estimation theoretic algorithms to
 - Determine both the magnitude and distribution of these emissions
 - Assess the impact of these emissions on the distribution of ozone
 - Investigate the processes linking emission source to ozone variability and distribution

QuickTime™ and a
TIFF (Uncompressed) decompressor
are needed to see this picture.

TES nadir provide a unique capability to help understand these processes by providing simultaneous profile retrievals of both ozone and carbon monoxide along with temperature, water vapor, surface, and effective cloud properties.

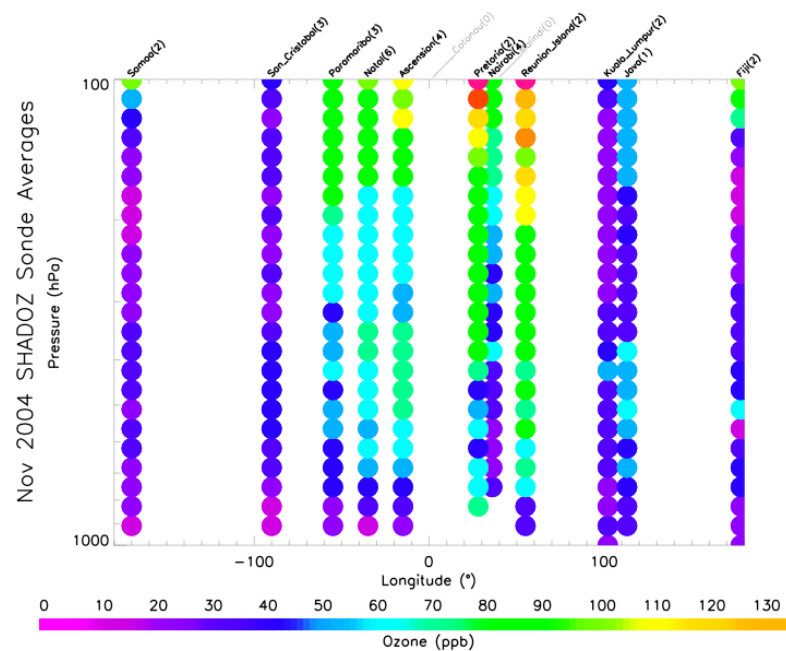
- For this study, 6 global surveys over the course of 12 days were used.
- The global survey mode for November 2004 produced 1152 observations per day spaced 5 degrees apart
- TES nadir foot print from the combined 16 detectors is 5.3×8.4 km
- V002 data was used for ozone and V001 data was used for CO

Comparison to ozone with the Shadoz network

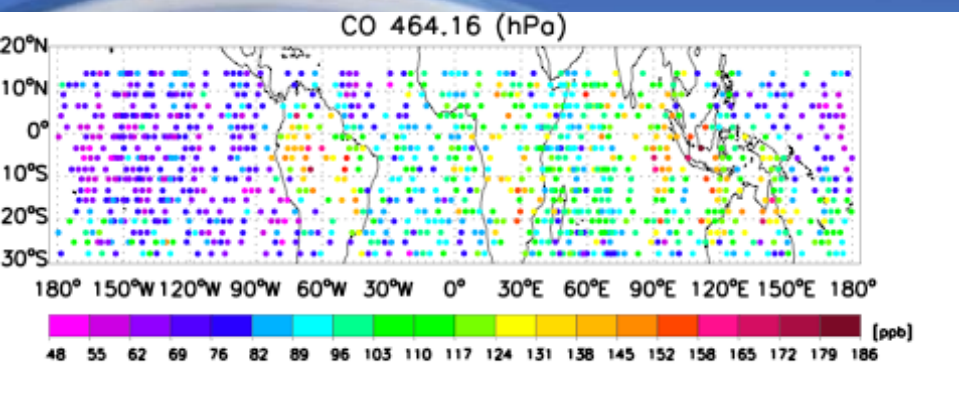


Shadoz mean of ozone was calculated from 1998-2004 for the month of November based on 11 sites covering a latitudinal and longitudinal range of 21S to 3N

Averages for November of 2004 are also included



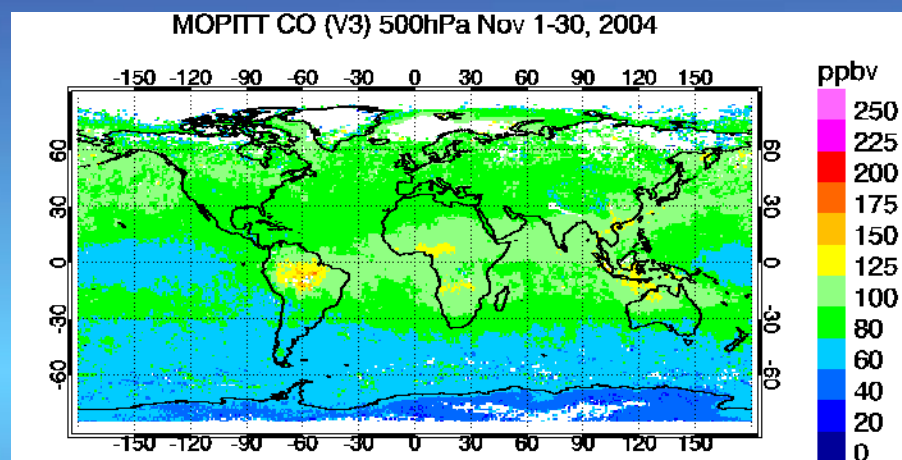
Comparison to MOPITT and AIRS tropospheric CO



Sciamachy near-infrared CO tropospheric column estimates

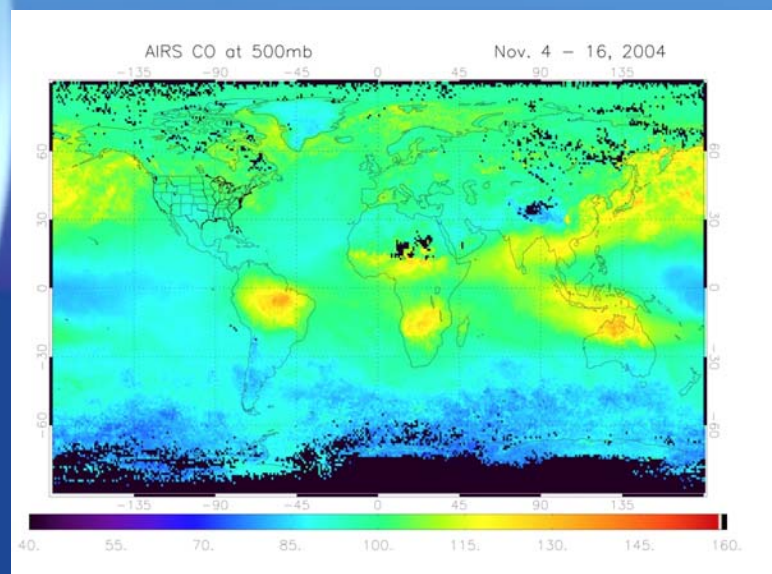
Elevated CO is observed between TES, AIRS, and MOPITT over South America, Africa, and Australia/Indonesia

A CO “plume” is observed by TES, AIRS, and MOPITT from the east of Australia extending across the south Pacific



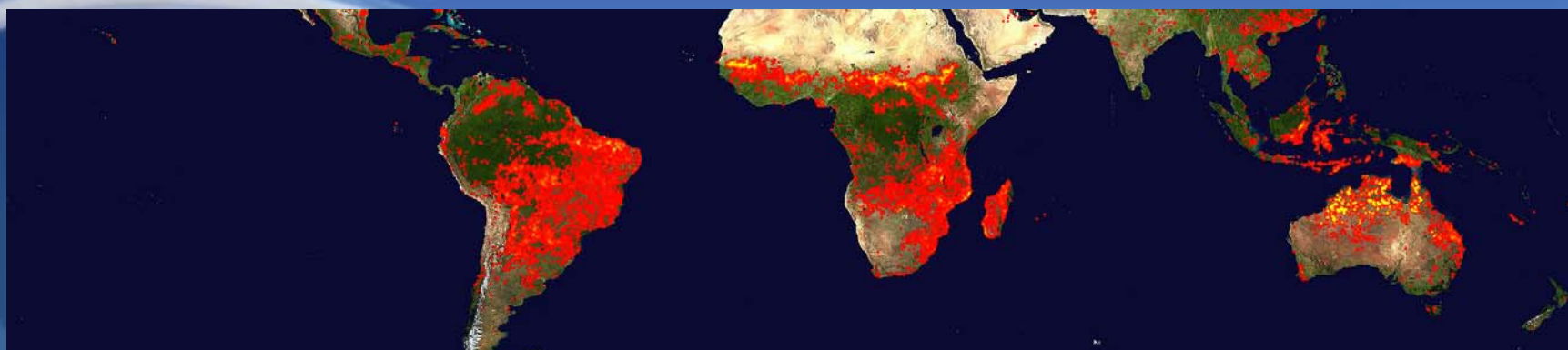
Gridded at 1x1deg from MOP02-200411??-L2V5*.hdf (apriori fraction < 50%)

MOPITT maps obtained from
http://www.eos.ucar.edu/mopitt/data/plots/mapsv3_mon.html



Signatures of biomass burning

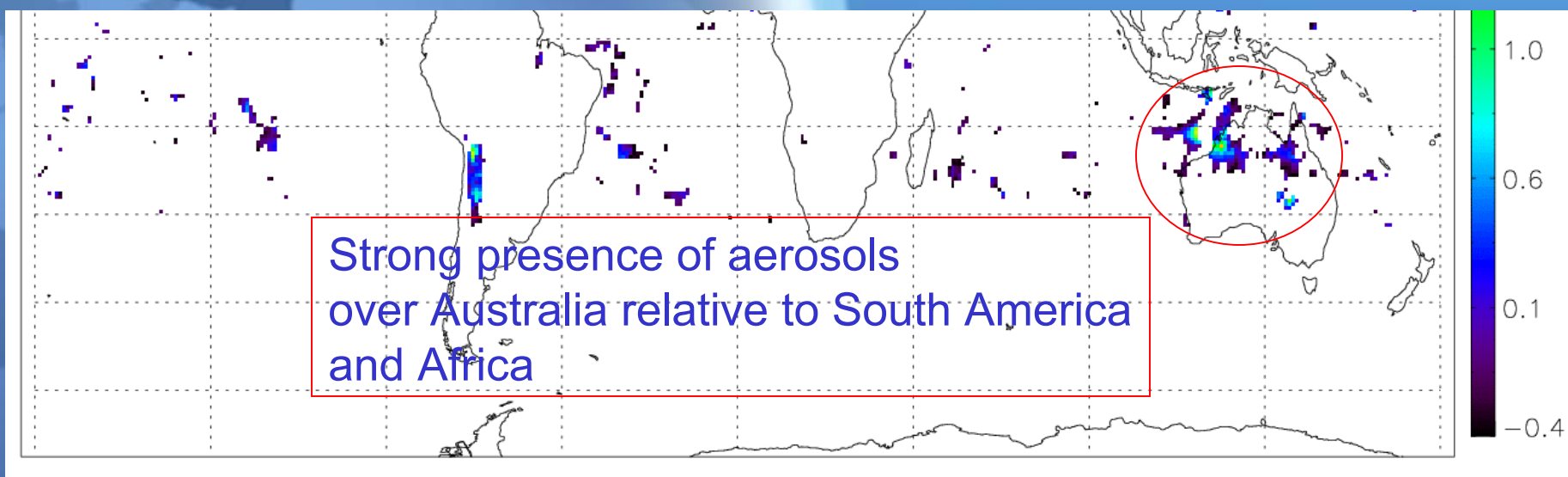
MODIS fire counts November 5-15



Red indicates at least one fire count
Yellow indicates a large number of fire counts

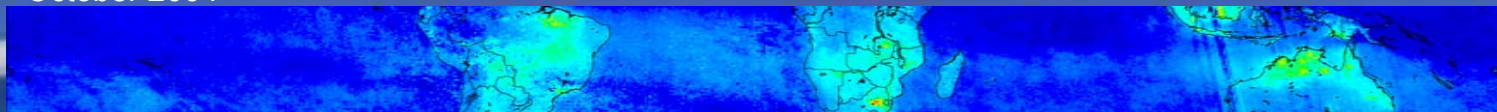
MODIS map obtained from
<http://rapidfire.sci.gsfc.nasa.gov/firemaps/?2005211-2005220>

Sciamachy absorbing aerosol index (AAI) November 2004

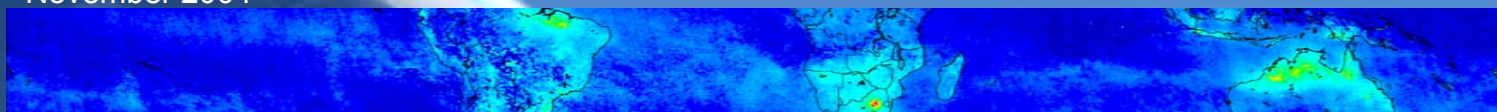


NO_x signatures and distributions

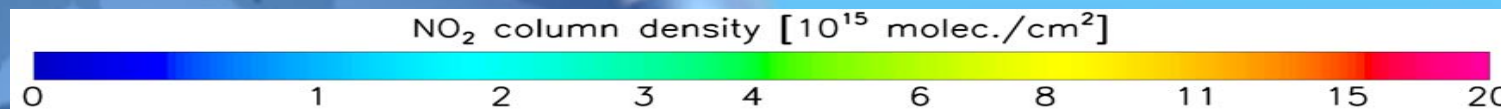
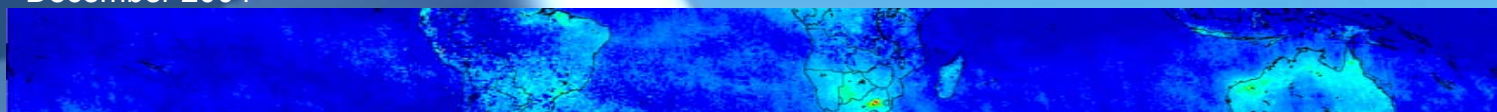
October 2004



November 2004



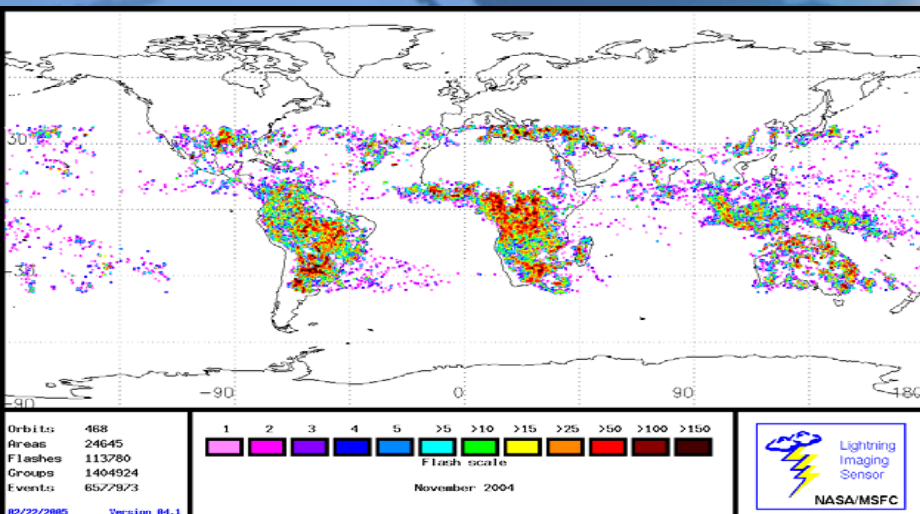
December 2004



<http://thunder.msfc.nasa.gov/data/query/distributions.html>

Lightning flash counts are high over South America, Africa, and Indonesia/Australia

OMI data indicates particularly high NO₂ columns in Indonesia-Australia relative sub-equatorial Africa and South America

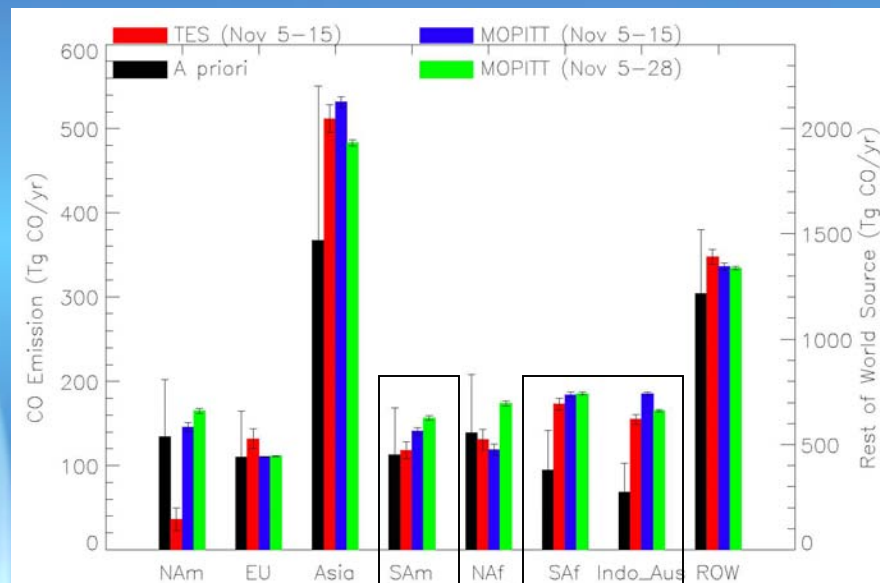


The magnitude and distribution of surface emissions can be estimate by assimilating TES CO into GEOS-Chem through a *Maximum a posteriori* (MAP) algorithm

EU
North America
North Africa
Asia
South America
South Africa
Indo/Aust
CHEM

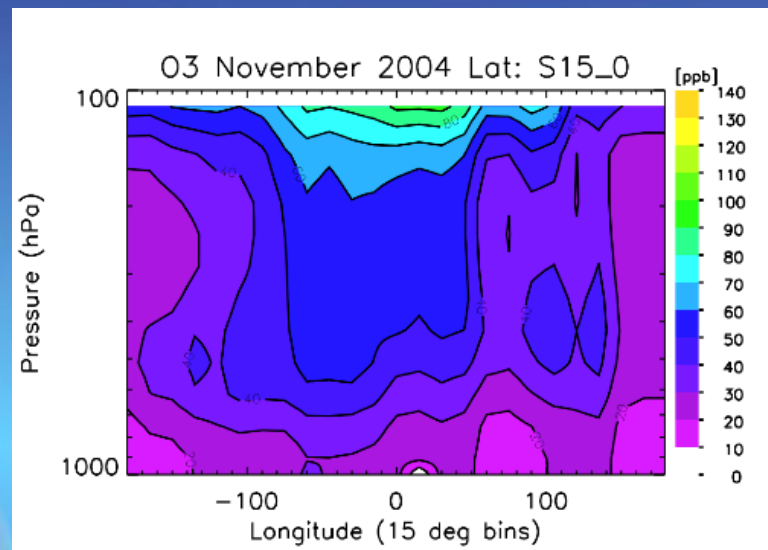
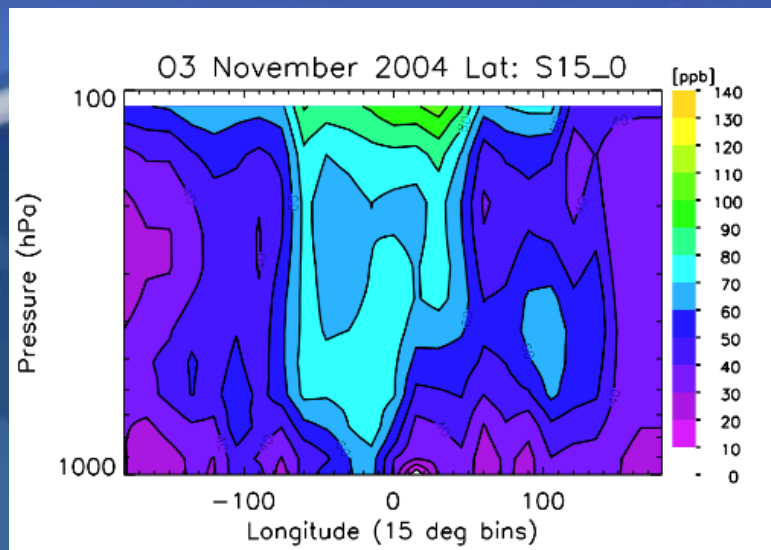
QuickTime™ and
TIFF (LZW) decompressor
are needed to see this picture.

- All sources include contributions from oxidation of VOCs
- Biomass, biofuel, and fossil fuel emissions are aggregated together
- CHEM includes methane oxidation and biogenic sources
- Forward model is GEOS-Chem, with specified OH
- Assumed an a priori source uncertainty of 50% for all sources; for the background chemical source we assumed 25% uncertainty
- Specified a uniform model error of 20%



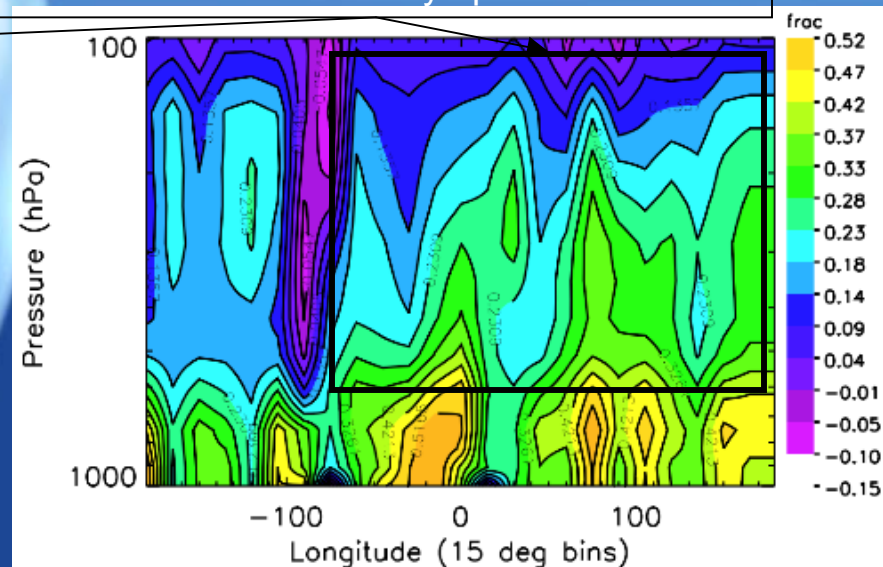
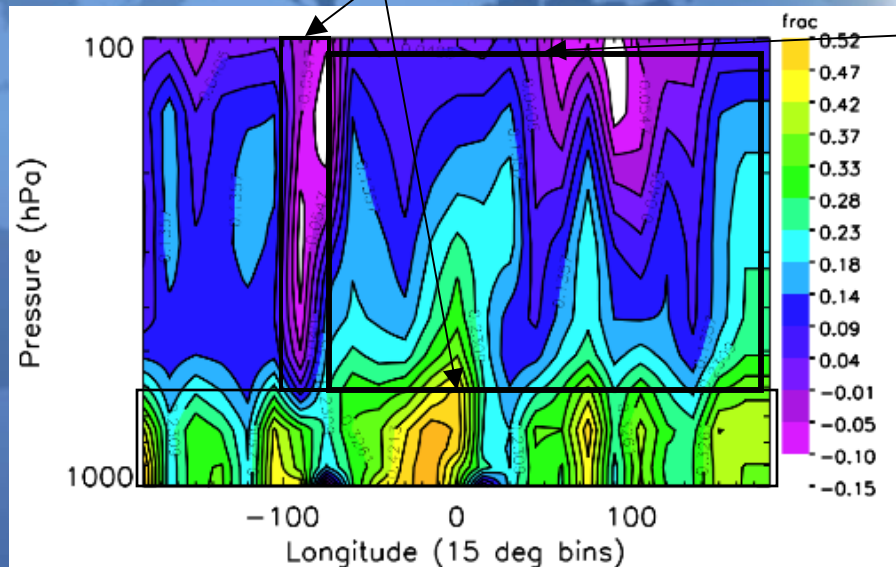
Estimates of surface emissions are a factor of 2 higher than the climatological inventory

Comparison of GEOS-Chem O₃ with a posterior emissions estimates and TES observations



Significant areas of disagreement below 500 mb and above South America

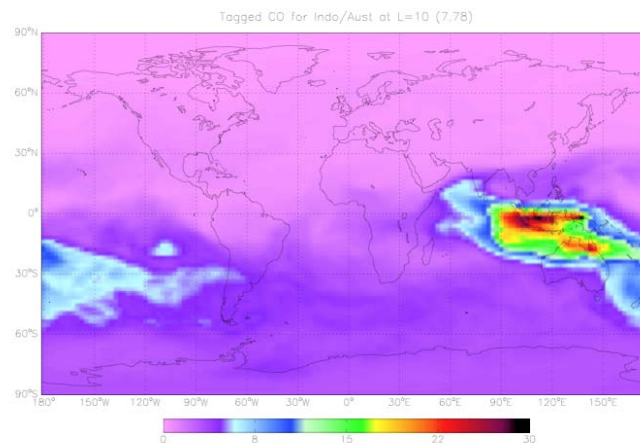
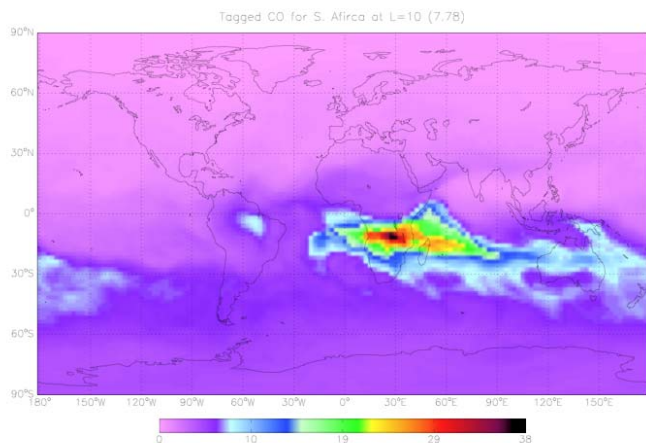
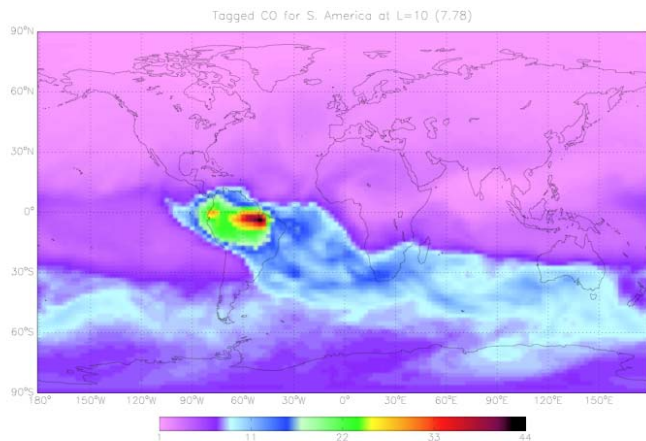
Substantial reduction in fractional difference between sub-equatorial Africa and Indonesia-Australia above 500 mb by up to a factor of 3



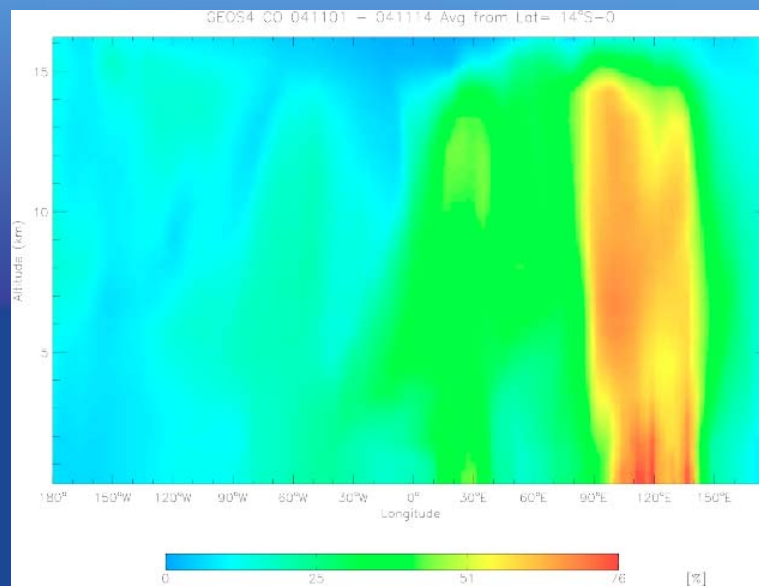
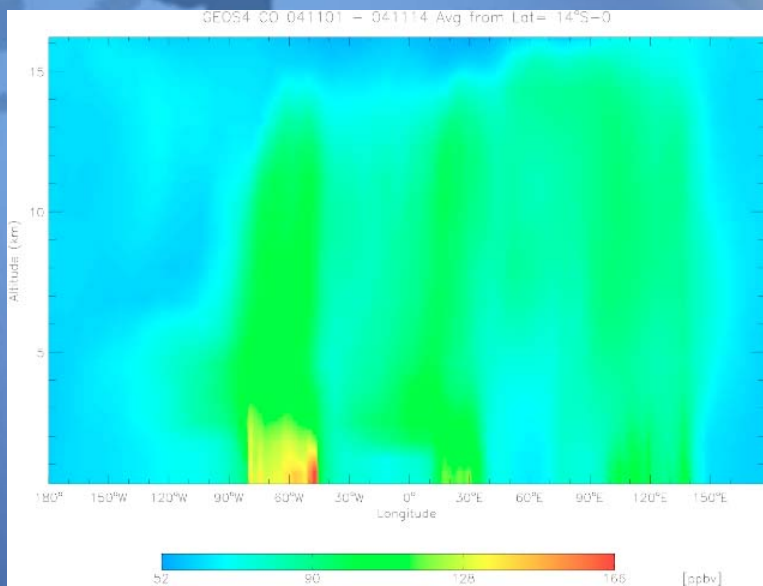
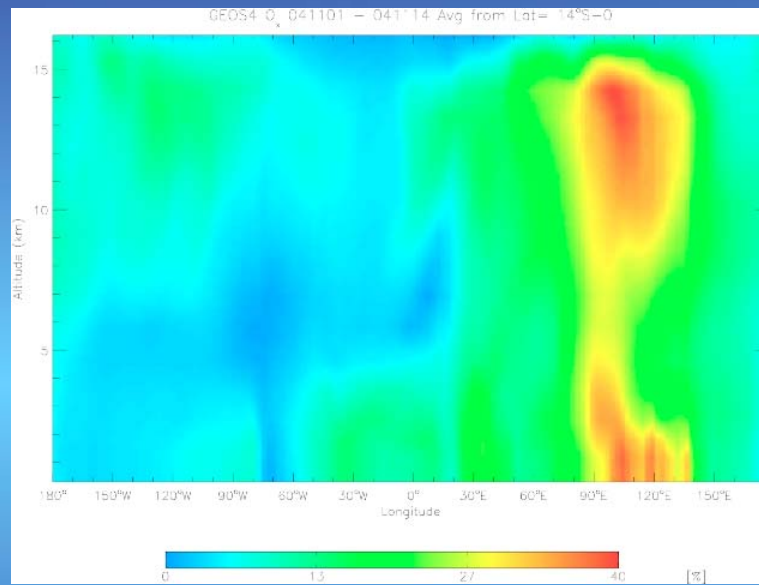
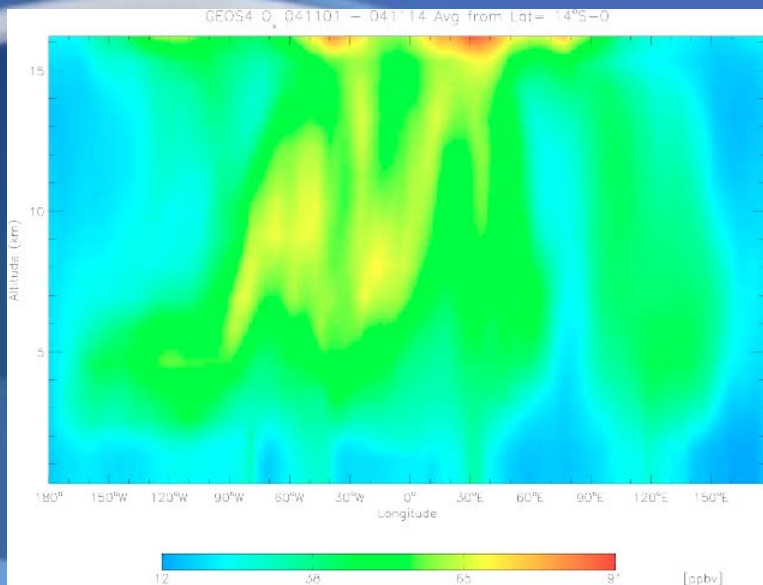
Process analysis

Tools used to estimate surface emissions can be important diagnostics to understand the processes controlling ozone formation

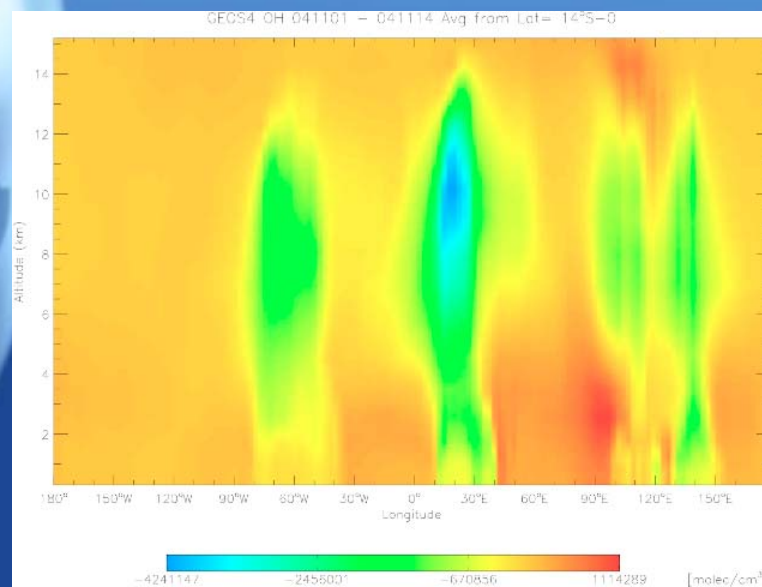
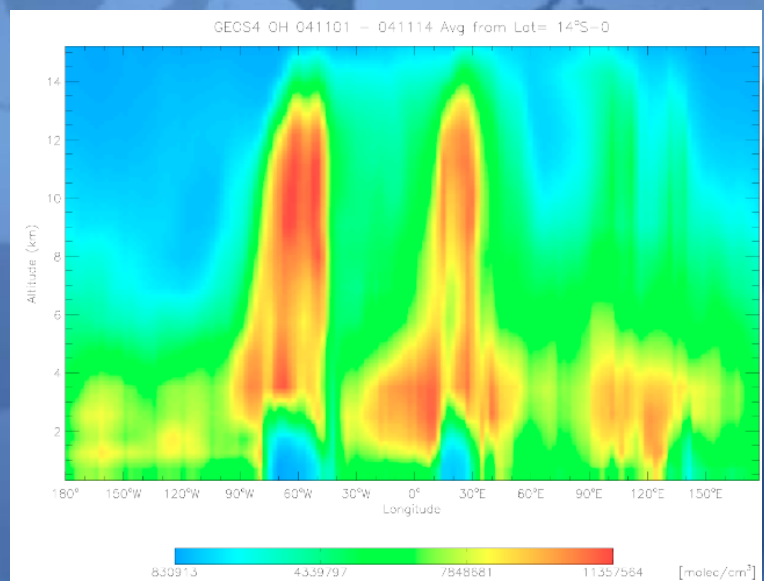
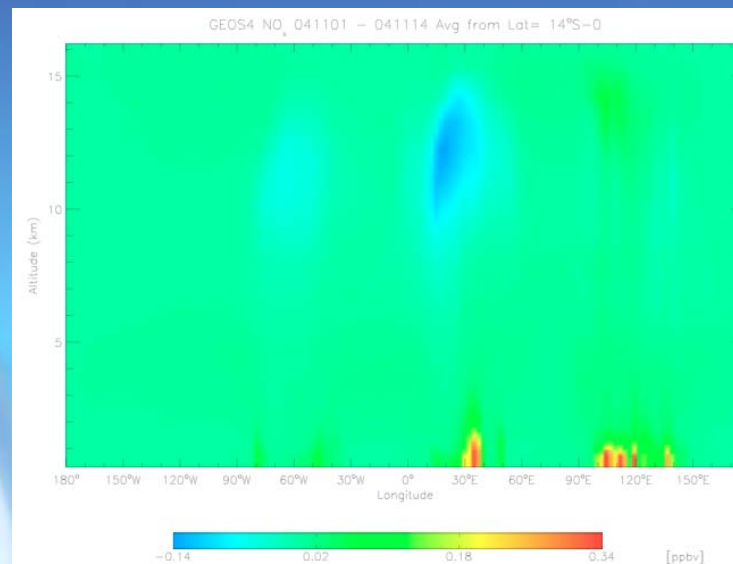
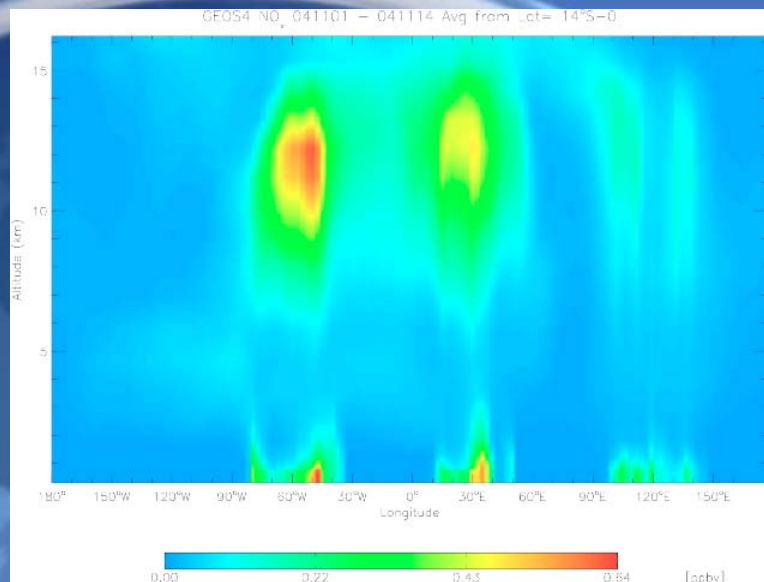
Tagged CO tracers from the 3 primary southern hemispheric continents indicate that the impact of surface emissions on ozone are fairly well separated



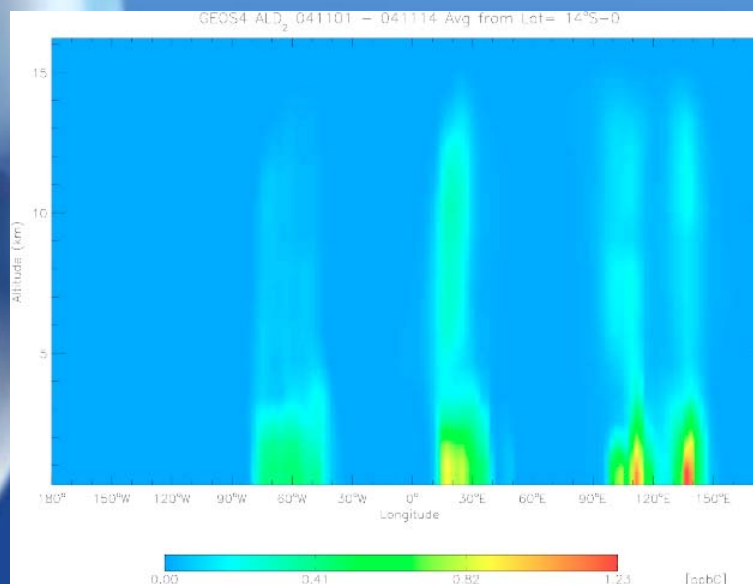
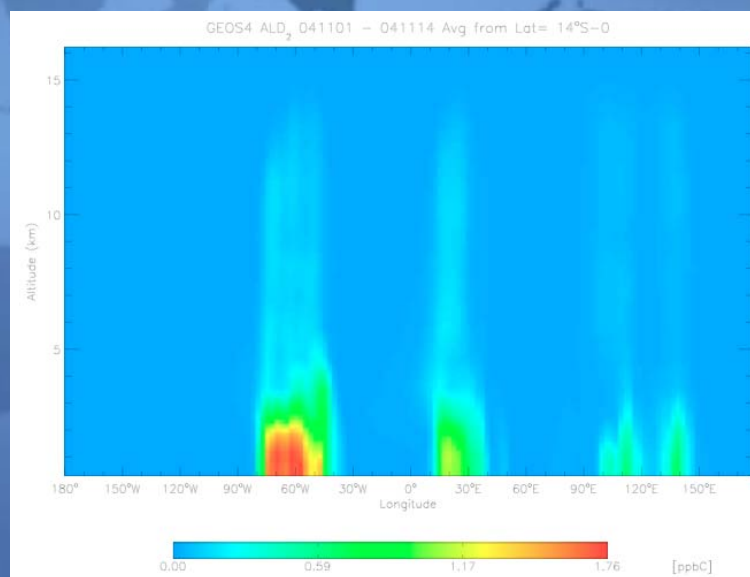
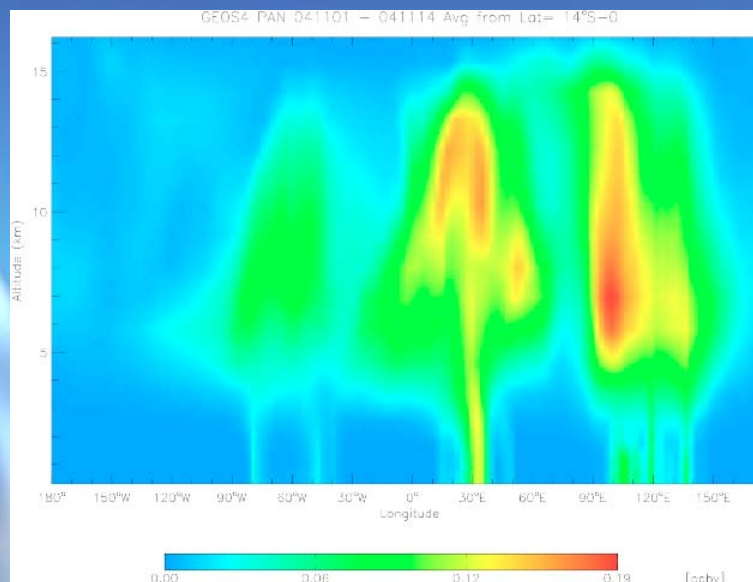
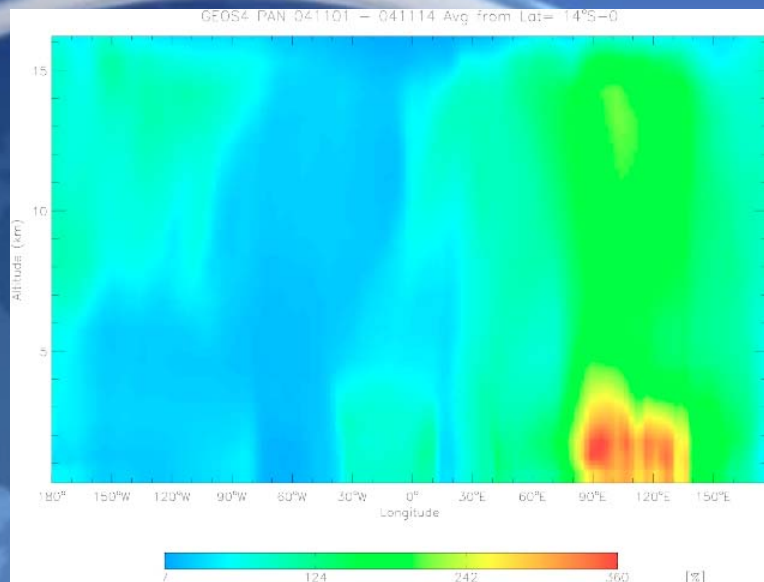
Sensitivity analysis for ozone and CO



Mediating species: NO_x and OH



Putting PAN in its place



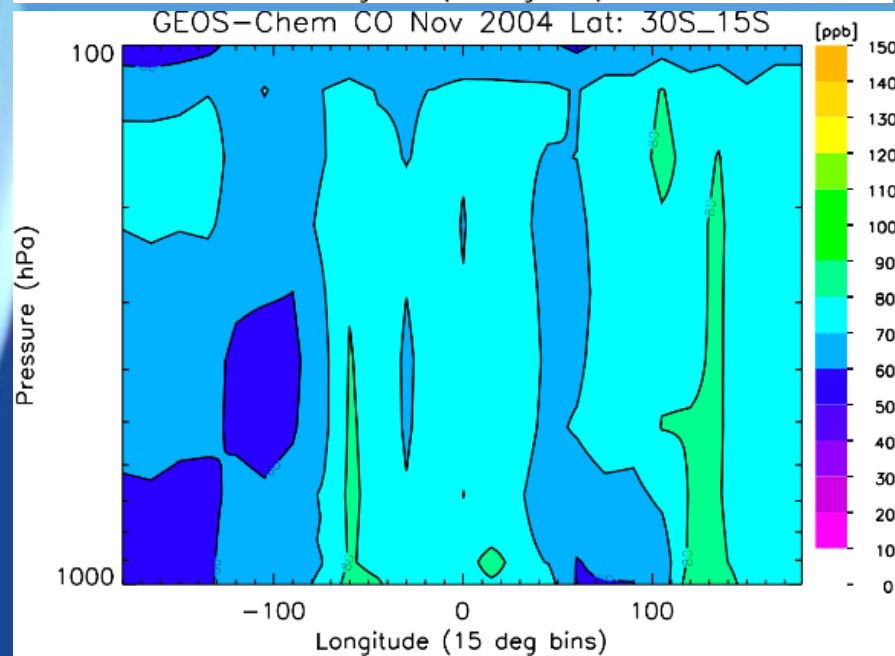
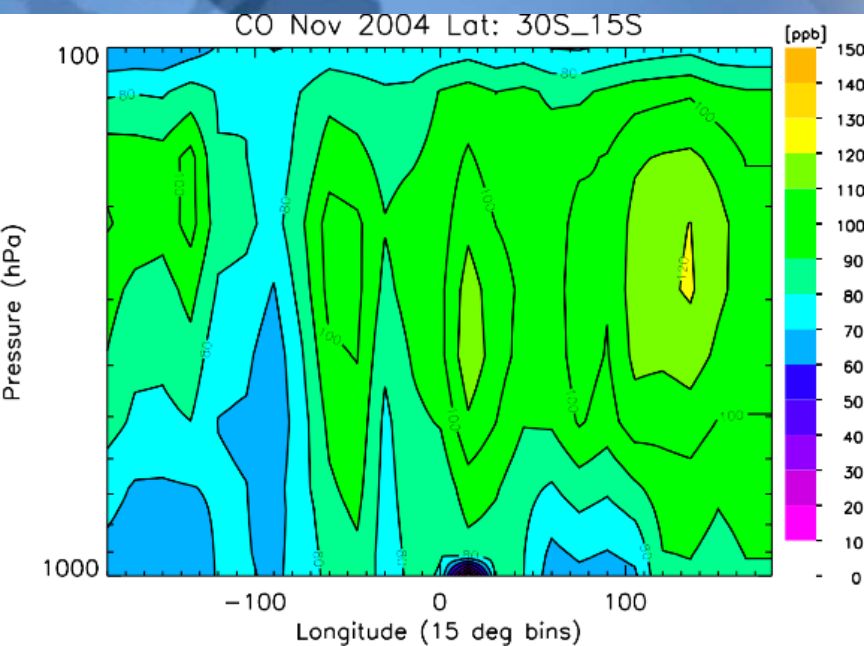
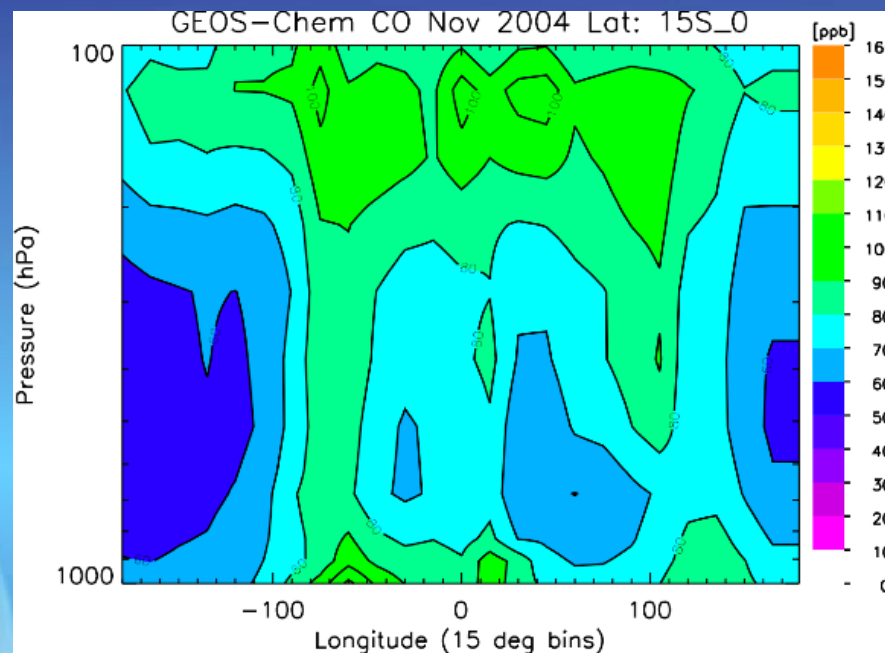
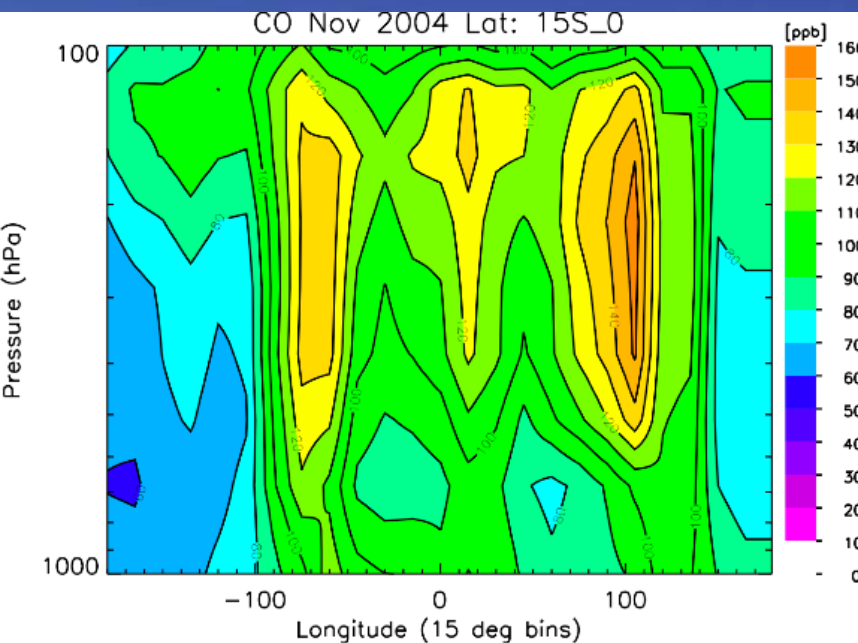
Conclusions

- TES retrievals show elevated ozone over the Atlantic--consistent with the "wave-one" pattern seen in the Shadoz network--as well as high ozone levels in Australia and Indonesia
- TES retrievals indicate elevated CO values over Australia/Indonesia are comparable to South America and Africa, consistent with AIRS and MOPITT CO estimates
- MODIS fire counts, Sciamachy aerosol and OMI NO₂ column measurements point to biomass burning as a source
- LIS flash counts suggest that lightning is also an important contributor of NO_x, particularly over South America and Africa
- Estimates of surface emissions from TES CO observations are up to a factor of 2 higher than climatological inventories
- Differences between TES ozone observations and *a posteriori* GEOS-Chem runs are reduced by up to a factor of 3, which suggest that surface emissions are an important contributor to ozone, particularly over Indonesia/Australia
- Sensitivity analysis can be used to investigate the key mediating processes that control ozone formation
- Advanced assimilation algorithms, such as 4DVar, adjoint analysis, Ensemble Kalman smoothing techniques will be critical for advancing the state-of-the-science in understanding the processes relating anthropogenic surface emissions to ozone variability

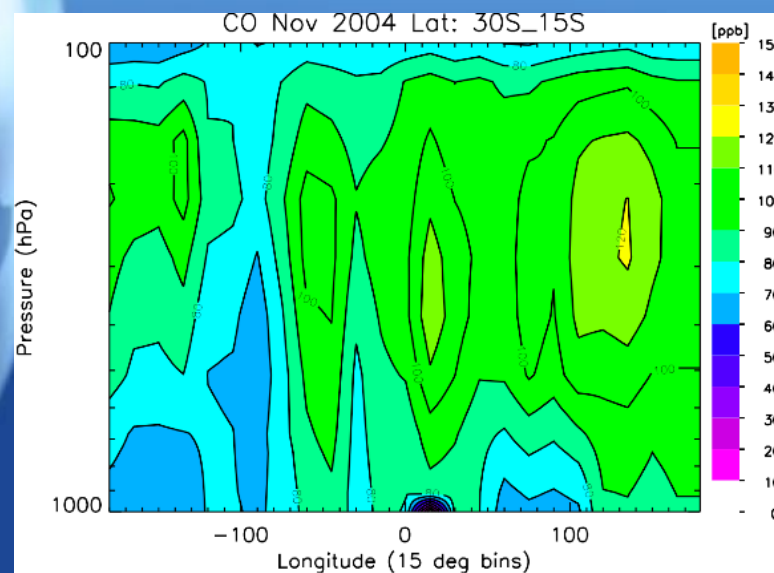
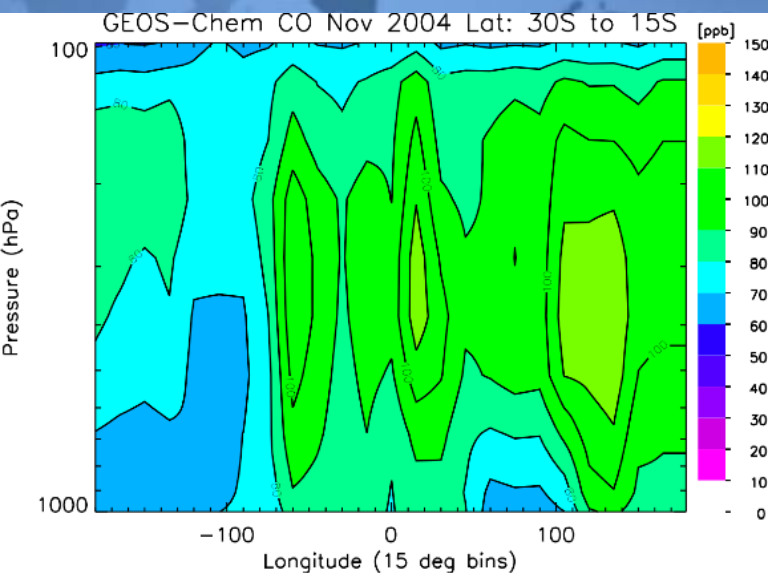
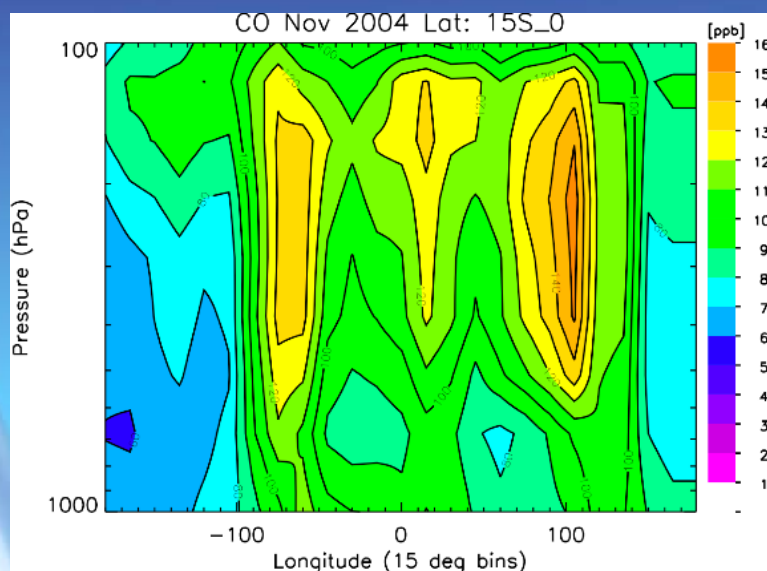
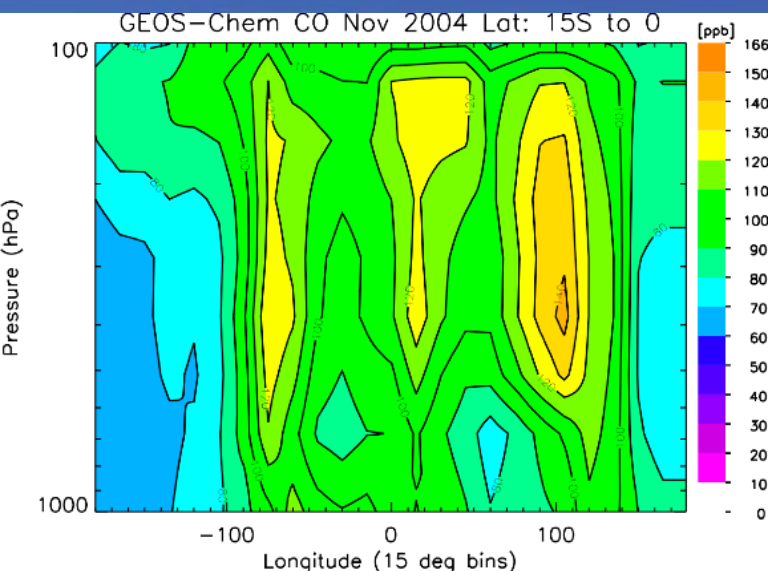
Backup slides



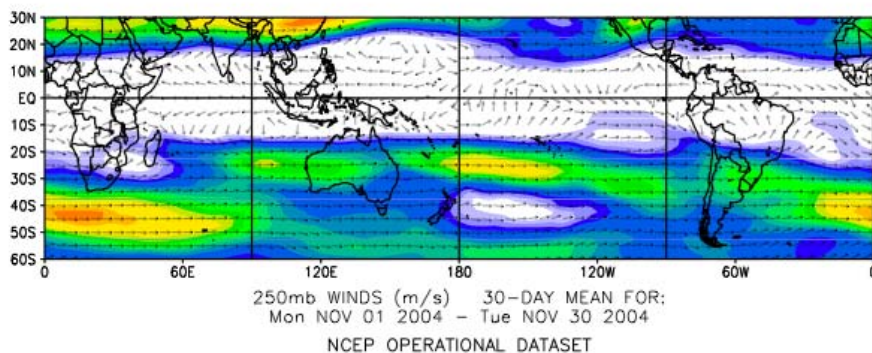
Comparisons to GEOS-Chem CO



Comparison of TES to posterior estimates



The distribution of tropical tropospheric ozone is governed by the complex interplay of chemistry and dynamics of both anthropogenic and natural origins:



Global transport
through advection

Convective lofting
of ozone and precursors

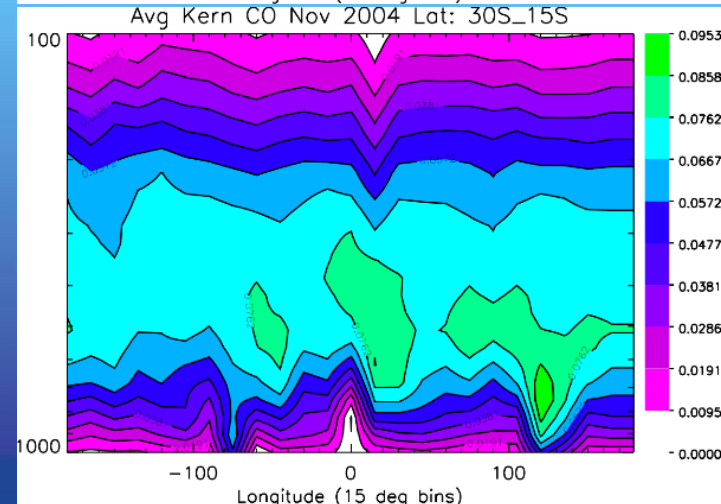
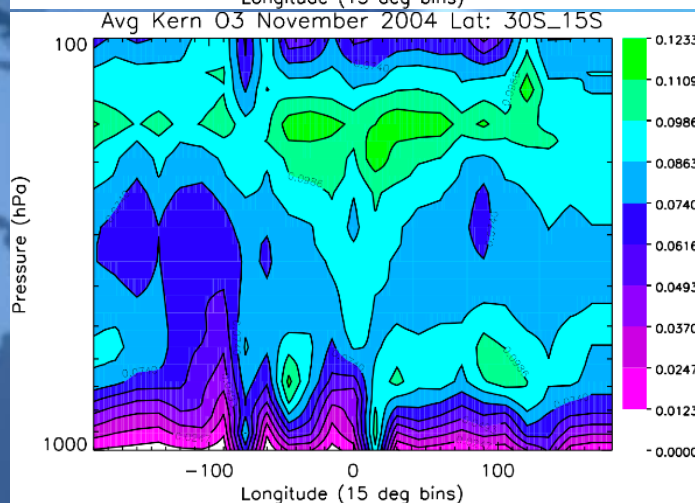
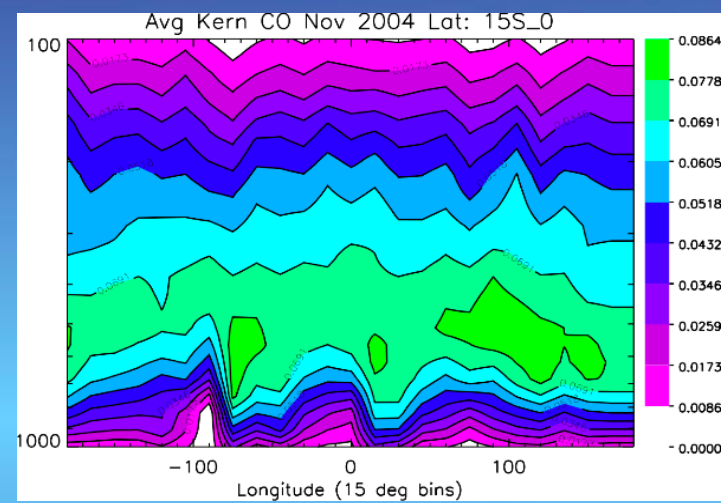
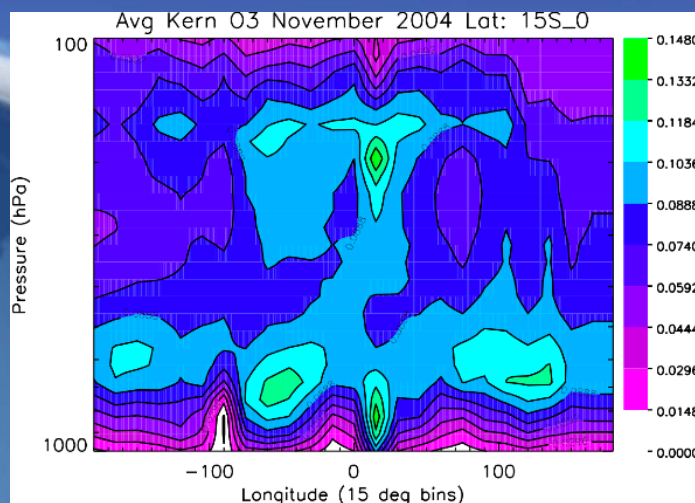
Lightning emitted NO_x

CO , NO_x , hydrocarbons, etc.
are produced at the surface
from biomass burning, fossil and
bio-fuel emissions



QuickTime™ and a
TIFF (Uncompressed) decompressor
are needed to see this picture.

TES ozone and CO sensitivity

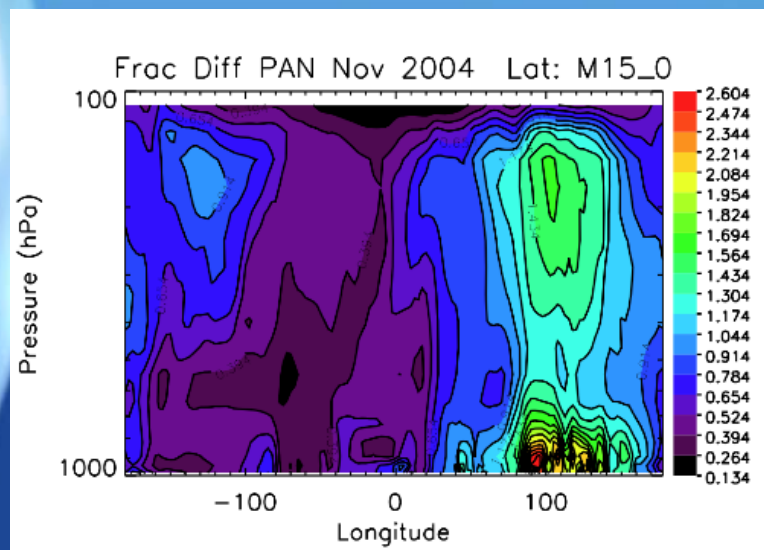
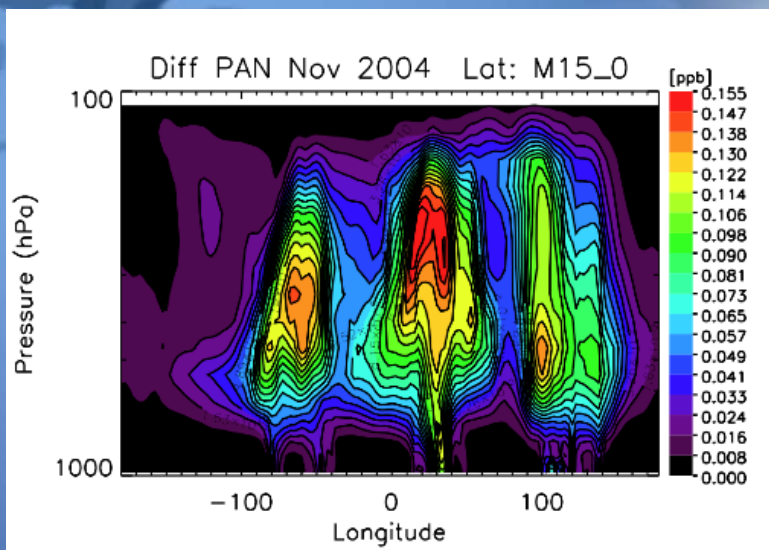
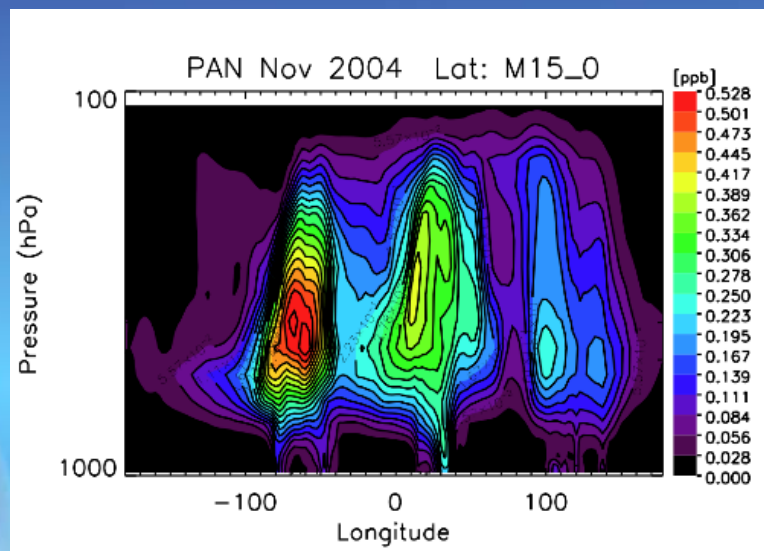
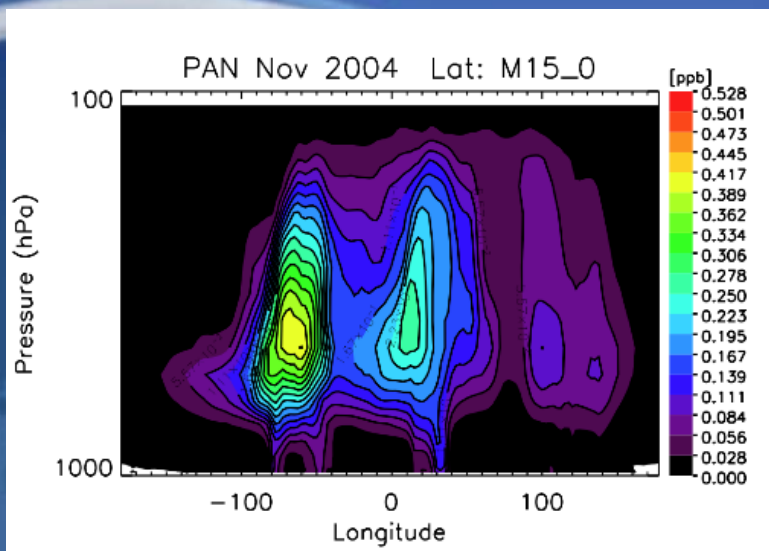


The averaging kernel is a measure of the sensitivity of the retrieval to the true state and metric for information gain

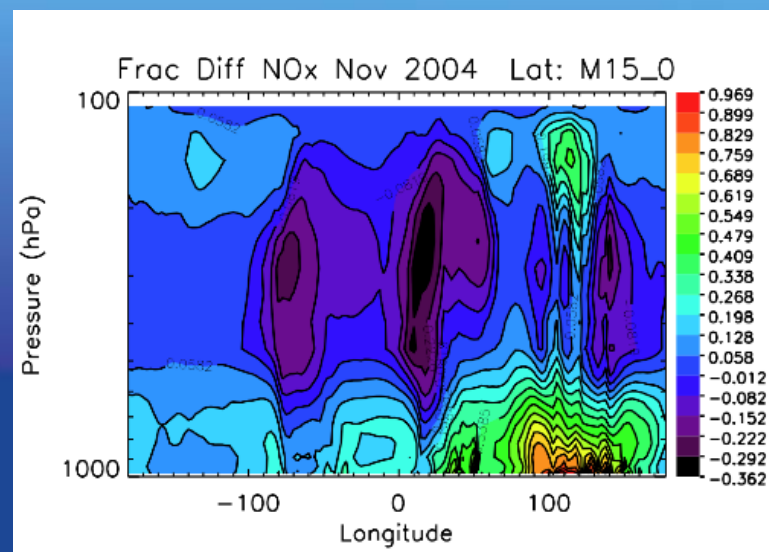
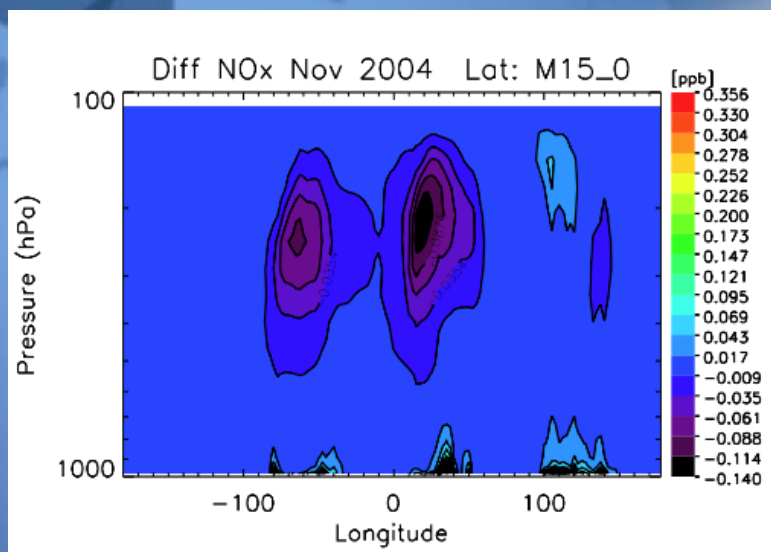
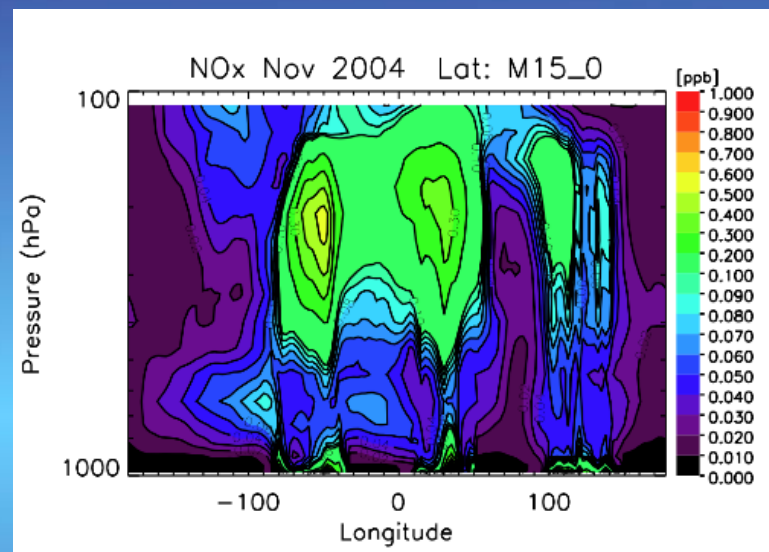
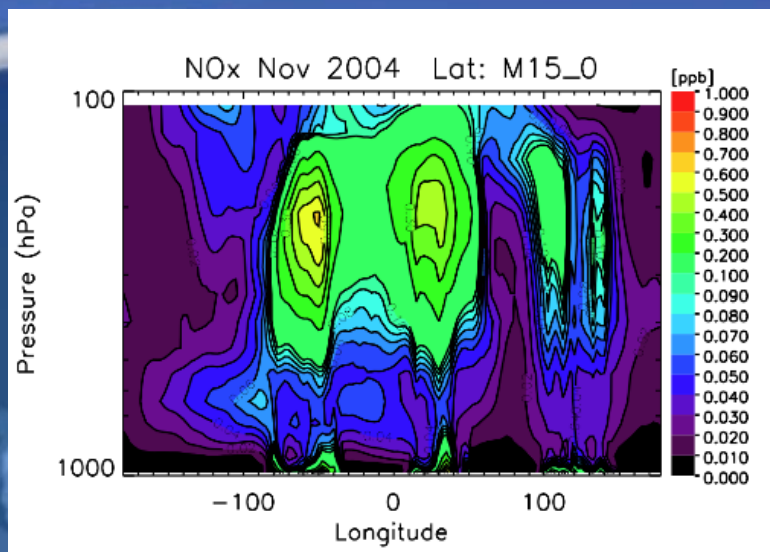
$$\mathbf{A} = \frac{\partial \hat{\mathbf{x}}}{\partial \mathbf{x}}$$

$$\mathbf{I} - \mathbf{A} = \mathbf{S}_{\tilde{\mathbf{x}}} \mathbf{S}_a^{-1}$$

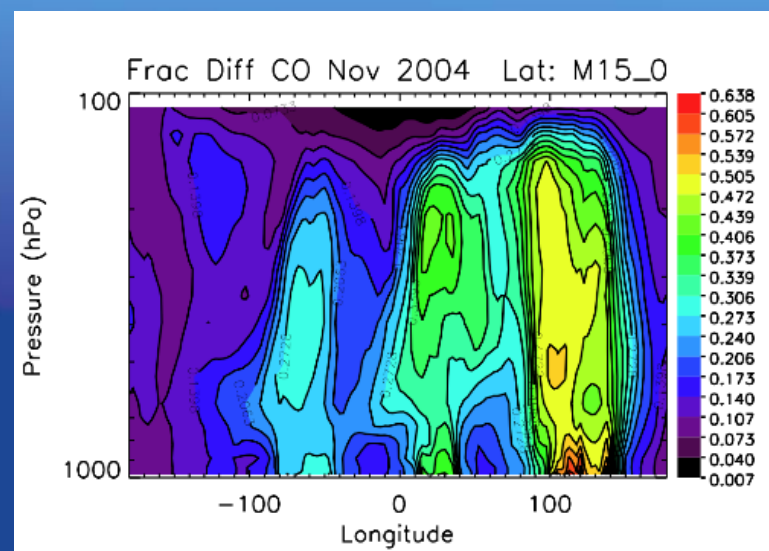
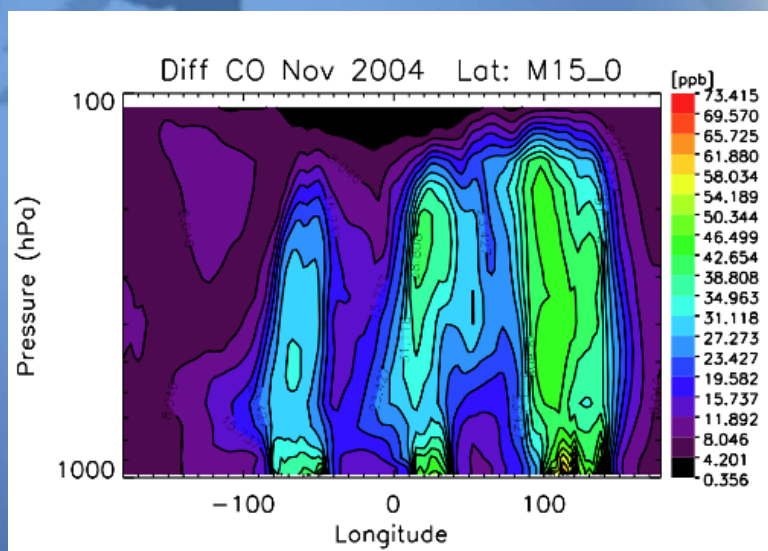
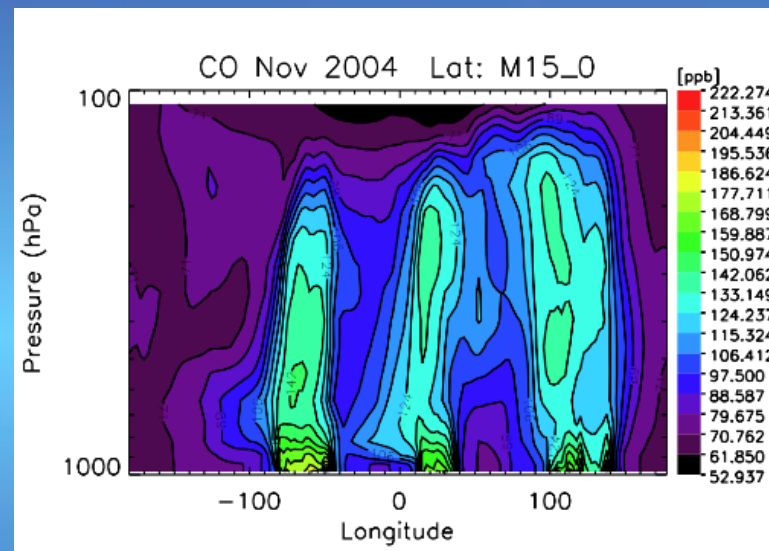
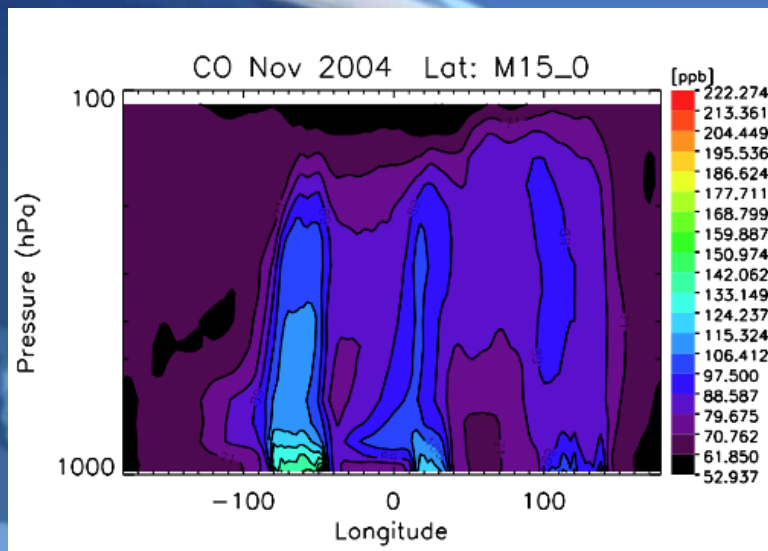
Sensitivity analysis: PAN



Sensitivity analysis: NO_x



Sensitivity analysis: CO



Sensitivity analysis: ozone

Prior

Posterior

